# Buffalo Lake

# Buffalo Lake Enhancement Project Summary Report





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# **SECTION I**

# Buffalo Lake Enhancement Project Hydrologic Study

(Conducted by Emmons & Olivier Resources, Inc.)

Prepared by: EOR

For the Buffalo Lake Protection and Rehabilitation District and Cason Land & Water Management, LLC

## Buffalo Lake Enhancement Project Hydrologic Study





April 11, 2025

#### Cover Images

Buffalo Lake Dam (November 20, 2025)

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#### 1. BACKGROUND

#### 1.1. Purpose and Scope

The Buffalo Lake District is evaluating a potential change in operation of the Buffalo Lake Dam (also known locally as the "Montello Dam") to raise the water level of the lake during parts of the year to improve conditions for lake users. The dam is owned and operated by the Wisconsin Department of Natural Resource (WDNR). The DNR awarded the District with a Surface Water Grant to study potential benefits and impacts of the proposed changes to the environment and landowners near the lake.

The dam has 4 operable gates plus an overflow spillway. Before 2019, they were not operated, and the lake level fluctuated based on inflows and the capacity of the overflow spillway. Starting in 2019, the DNR began to operate the gates to maintain the level of Buffalo Lake consistent with the water levels specified in the 1976 operating order. The operating order specifies a winter maximum lake level corresponding to 8.0 ft on the staff gage mounted on the dam and a summer maximum of 8.5 ft. The order specifies that the summer maximum applies to the period from May 20 through October 1.

No change in either the summer or winter maximum stages is currently proposed. Rather, the District proposes to increase the duration of the summer maximum to May 1 – October 15, an additional 20 days in May and an additional 14 days in October.

This study evaluated potential impacts that the proposed dam operation change could have on the following issues:

- Peak flood elevations on the lake
- Groundwater levels around the lake
- Water temperature in the Fox River downstream of the dam
- Lakeshore erosion
- Wetland community composition and quality

#### 2. DAM OPERATION AND LAKE LEVEL DATA

#### 2.1. Lake Level Records

The WDNR provided a spreadsheet with data on the lake stage and dam gate operation for the period of June 3, 2019 through July 2, 2024 (Uriah Monday, written communication, 2024). WDNR photographs the staff gage on each visit and records the lake level in the spreadsheet. A staff gage reading of 8.0 ft corresponds to elevation 769.0 ft in the National Geodetic Vertical Datum of 1929 (NGVD29) according to WDNR records. The Flood Insurance Study (FIS) lists a vertical datum conversion between NGVD29 and the more recent North American Vertical Datum of 1988 (NAVD88) of NGVD29 – 0.1 ft = NAVD88. Therefore, a gage reading of 8.0 ft equals 768.9 ft NAVD88 (**Table 1**). Unless otherwise noted, this report describes elevations in NAVD88.

Gage Reading (ft)	NAVD88 Elevation (ft)	
8.0	768.9	
8.5	769.4	
9.0	769.0	

Table 1. Conversion of staff gage measurements and NAVD88 elevations.

Add 760.9 ft to gage readings to calculate NAVD88 elevations.

#### 2.2. Dam Operation

The dam has 3 sluice gates and one split leaf gate, all with a sill elevation of 761.5 ft. The dam's auxiliary spillway is a weir with a 168-ft-long crest at elevation 768.65 ft (7.75 ft on the local gage datum). There is additionally a fish passage with an invert of 768.2 ft (7.3 ft on the local gage datum) and a flow width of 20'.

The gates have typically been checked on weekly visits by WDNR staff and adjusted as needed to comply with the water level order. It appears that the gate operation is based on professional judgement of the operator and not a written operation plan. The gates have been partially open by variable amounts most of the time during the winter water level period and closed most of the time during the summer season, with the gates open occasionally to pass high flows.

Spring operation records for 2020 – 2024 indicate that, after the gate settings are adjusted on May 20, it takes about a week for the water level to approximately reach the summer maximum. In 2020, the lake was already above the summer maximum at 8.8 ft on May 20, so gates were opened to lower the water level. Conditions in 2023 were very dry, and the lake could not be raised all the way to 8.5 ft for most of the summer.

In late summer, dry conditions typically have made it difficult to maintain the lake at 8.5 ft, so that little or no gate adjustment has been needed on October 1 to drop the lake to the winter maximum of 8.0 ft.

#### 2.3. Lake Inflows

The WDNR also provided estimates of daily lake inflows for October 1<sup>st</sup>, 1991 through July 9<sup>th</sup>, 2024, and weekly outflows for June 3<sup>rd</sup> 2019 through July 2<sup>nd</sup> 2024. Inflows into Buffalo Lake are based on flows at the USGS gage on the Fox River downstream of the lake at Berlin prorated by drainage area. Lake outflows are estimated by DNR's hydraulic calculations of the discharge through the dam based on the gate settings and overflow spillway capacity. Comparison of lake inflows estimated by the gage transfer method and lake outflows estimated with hydraulic calculations for the dam gates shows that these methods estimate similar low flows, but high flow estimates for the dam gates are much higher. Recession curves after high flows are similar for both methods (**Figure 1**). This provides confidence that the estimate lake inflow and outflows are reasonable.



Figure 1. WDNR estimates of Buffalo Lake inflows (blue) and outflows (orange).

#### 2.4. Shoreline Changes

No elevation data is available to precisely define the lake's shoreline position at the ordered levels of 8.0 ft or 8.5 ft. The most recent topographic data for Buffalo Lake and its shoreline is the Marquette County LiDAR survey from 2018. At that time, the lake stage was 770.36 ft (9.46 ft on the gage), which is above both the summer maximum. A bathymetric map from WDNR for 1967 (**Figure 2**) illustrates the shoreline at a stage of 8.0 ft and depth contours below the water level. However, this map obviously does not use modern terrain data and its scale is too small for precise comparisons with other data sources.

Information on how changes in lake stage in the range of 8.0 ft to 8.5 ft affect the shoreline position can be gleaned from aerial photographs from three recent years on dates close to stage measurements by the WDNR. On March 11, 2021, the lake stage was between 8.14 ft and 8.22 ft. On June 30, the stage was between 8.5 ft and 8.64 ft, and on May 4, 2024 the lake was between 7.93 ft and 8.03 ft. Images from these dates are compared for several locations in **Appendix A**. These images illustrate that the lake inundation extents and surface area vary little across the operational range.



Figure 2. A portion of the 1967 WDNR bathymetric map.

#### 3. LAKE FLOOD LEVELS

Two related but distinct issues were evaluated in this analysis of flood levels on Buffalo Lake. One is whether the proposed change to the dam operation would affect the effective regulatory Base Flood Elevation (BFE) for the 100-year event listed in the Flood Insurance Study (FIS) for Marquette County. This depends on how the flood elevation was calculated in the FIS. The second issue is whether and by how much flood levels would actually change if a flood occurred during the periods in May or October when the lake is proposed to be raised from 8.0 ft to 8.5 ft.

#### 3.1. Flood Insurance Study Review

Flood floods into Buffalo Lake used in the Flood Insurance Study (FIS) are based on the hydrologic analysis in the Columbia County FIS, because flood flows in the Fox River upstream of the lake are affected by overflow from the Wisconsin River into the Fox River at Portage, Wisconsin. The FIS used a coincident frequency analysis of the Wisconsin and Fox Rivers and an estimate of overflows between the watersheds.

Hydrologic and hydraulic flood calculations for Buffalo Lake were performed by the U.S. Army Corps of Engineers (USACE) in 1999 using a UNET model for the Fox River, the results of which were then used to initialize a HEC-2 model of flood elevations. EOR spoke with the USACE modeler, Terry Zien, for insights into that analysis and received written records from the USACE. The UNET model uses a lake stage of 769.37 ft NGVD 1929 at the dam as the initial condition before routing the upstream hydrograph through the lake. This corresponds to a gage reading of 8.37 ft, which is between the summer and winter maximum water

levels. Documentation from the USACE states that this starting elevation was selected using a rating curve for the dam developed in 1980 by Owen Ayers Associates. Presumably, the USACE selected a representative low flow in Buffalo Lake to calculate this elevation, but those details are not available. The UNET model calculated a peak 100-year discharge of 3829 cfs and a 100-year stage of 771.99 ft (NGVD 1929) at the Montello dam. This peak discharge is considerably lower than that for the Fox River upstream of the lake, indicating a substantial routing effect through the lake.

Because the UNET flood elevation at the Montello Dam was used as the downstream boundary condition for the HEC-2 model, either model could reasonably be said to be the source of the elevations at the Dam. Moreover, the downstream boundary for the UNET model is a rating curve referenced to the 1980 National Dam Safety Program Inspection Report, Fox River, Montello Dam (Inventory No 01015). We were unable to locate a copy of this document, however the rating curve was obtained from the UNET input files.

Comparing the rating curve in the UNET model to model results at the same discharges for a hydrologic model developed by EOR using the HEC-HMS software (described below in Section 3.2.3) shows a near-identical stage-discharge relationship when the dam gates are fully closed; for this reason, we conclude that the FIS flood elevations were developed with the assumption that the dam gates were fully closed.

#### 3.2. Analysis

The proposed dam operation change could theoretically affect the lake flood stage if the flood occurred during the period when the lake level would be at 8.5 ft instead of 8.0 ft. EOR evaluated the impact of the proposed water level change on lake flood stage by 3 methods: evaluating the seasonality of historical floods, a comparison of flood hydrograph volume with lake storage, and performing hydrologic modeling.

#### 3.2.1. Seasonality of Flood Flows

Stream flows and flood risk are non-uniform during the year, and so the timing of the proposed changes must be examined in the context of the seasonal variation of high flows.

According to an analysis performed by Uriah Monday of the WDNR, flows in the Fox River peak in the spring, roughly between mid-March and early June (**Figure 3**). Median flows during this period range from 500 to 600 cfs, well up from the 250-300 cfs median values for much of the remainder of the year. This suggests that the marginal risk of increased water levels is significantly concentrated during the spring period of interest, rather than the fall period.

Peak annual flows do not always coincide with periods of generally high flow. For the period 1992 – 2021 there were 8 recorded years with annual peaks higher than 1000 cfs; of these, half took place during the summer and early fall, and only one took place during the periods of proposed change (1271 cfs on May 8, 2012), with the remainder occurring in March and April.

We conclude that the antecedent conditions that create increased vulnerability are likely concentrated during the spring period of interest, while the individual high flows that would be likely to trigger an extreme flood event are more broadly distributed throughout the summer months.



Figure 3. Exceedance probability flows for WY 1992 – 2022 (from WDNR with annotation of proposed spring and fall extension of summer maximum lake stage added by EOR).

#### 3.2.2. Volumetric Comparison

The 100-year flood hydrograph calculated for Buffalo Lake from the UNET model output has a full volume of 33,708 ac-ft. This includes a 1-day "warm up" period with a constant discharge of 300 cfs. The flood discharge peaks at the end of day 5 and is still falling after day 9 (the end of the modeled hydrograph). Considering only the rising limb of the hydrograph, from its start on model day 2 to its peak at the end of day 5 yields a volume of 13,737 ac-ft.

The incremental storage volume between the starting stage of 8.37 ft assumed in the FIS and the 8.5 ft maximum summer level can be calculated by multiplying the lake area (2179 ac per the WDNR) by the water level difference of 0.13 ft. This yields a storage difference of 283 ac-ft. This is 2% of the volume of the rising limb of the 100-year flood hydrograph, a small enough difference that the peak lake stage should be essentially the same for a 100-year event with the lake starting at either the elevation assumed in the FIS or the 8.5 ft summer maximum stage.

#### 3.2.3. Hydrologic Modeling

Hydrologic modeling was used to evaluate the difference in flood stage on the lake if a 100-year event were to occur during one of the periods when the lake stage is proposed to be raised from 8.0 ft to 8.5 ft. The USACE no longer supports the UNET model used in the FIS, and the FIS analysis did not simulate the dam gates in a way that is conducive to testing the impact of gate operation on water levels. Therefore, EOR built a hydrologic model using the HEC-HMS software to evaluate lake inflows, outflows and stage. This model was used to evaluate the effects of antecedent lake stage and gate operation on flood peak elevations.

HEC-HMS model construction details include the following.

• <u>*Reservoir*</u> – Limited bathymetry data were available, however over the very limited elevation range with active storage there is little variation in the marginal storage per unit of elevation change. The stage-area relationship is shown in **Table 2**.

Stage (ft)	Area (ac)	Source		
0	0	Per the 1967 WDNR map, the lake stage is defined from the lake invert		
8	2179	Normal stage lake area published by WDNR		
11.2	2900	100-year flood elevation, area measured from 100-year floodplain maps.		
12	2900	Conservative assumption of straight sided storage; not active in this model.		

able 2. Buffalo La	ke stage-area r	elationship use	ed in HE	C-HMS model

- <u>Dam</u> The model includes a single sluice gate representing the 4 gates with elevations dimensions taken from the dam plans and an opening width equal to the opening width of the 4 gates together. In reality, one gate is a split-leaf type, and the gates are operated independently; however the goal of this modeling was to represent the aggregate capacity of the gates accurately enough to evaluate impacts of the proposed water level change, not to simulate individual gate operation. The model also includes the fish ladder and auxiliary spillway weir.
- <u>Historical Flows</u> The model uses WDNR's lake inflows estimated based on drainage area ratio with the USGS gage at Berlin. For modeling efficiency, we simulated slightly more than 1-year period using the annual flow series including the 2008 water year, with an additional 3 months at the end (October 1, 2007 – December 31, 2008). This period contained the largest event during the period of record, which had an estimated peak inflow to Buffalo Lake of 1776 cfs plus an apparent snowmelt event of 1135 cfs.
- <u>FIS flows</u> We used the 9-day hydrograph from the UNET model developed by the USACE and used in the FIS (**Figure 4**). The peak discharge is 3829 cfs.



Figure 4. UNET discharge hydrograph. Horizontal axis units are days.

#### Historical Flow Data Simulation

A continuous simulation of historical flows was performed to evaluate the impact that gate settings have on the flood level of the lake. As noted above, we simulated a little over 1 year using 2007 and 2008 data for modeling efficiency. The model does not attempt to simulate actual weekly adjustments of the gates. Rather, a range of constant gate openings was modeled and compared to ascertain the difference that makes for lake flood stage. **Figure 5** shows model runs for the 4 gates being closed, and for openings of 0.5 ft, 1 ft, and 2 ft (all gates open the same amount). The runs most closely representing actual historical operation of the dam are the scenarios with the gates closed and gates open 0.5 ft. The lake stage during low flows for these scenarios is generally between 8.0 ft - 8.5 ft for the gates closed and between 7.5 ft - 8.0 ft for the 0.5-ft-opening scenario.

Also note that these simulations illustrate that gate operation can make a substantial difference in the lake flood elevation. For the largest event simulated in 2008, the peak lake stage is approximately 0.7 ft lower for the scenario with the gates open 2 ft vs. the gates-closed scenario. These means there is an opportunity to mitigate flood impacts through gate operation during high flow events.



Figure 5. HEC-HMS simulation of 2007-2008 flows for gates closed (upper left), gates open 0.5 ft (upper right), gates open 1.0 ft (lower left) and gates open 2.0 ft (lower right). Lake stage shown in light blue.

#### Impact on the Flood Insurance Study Elevations

The HEC-HMS model developed for this project was used to establish a rating curve based on the reconstructed dam geometry with the modeled gates fully closed, and this was compared to the published FIS values. Through the range of flows up to the 100-year flow (3829 cfs) the calculated water surface elevations at the dam match within 0.02 ft, with the exception of the highest tested flow of 4150 cfs where the two figures deviated by 0.07 ft (**Table 3**). Based on this near match, we conclude that the FIS elevations were developed assuming the dam gates were fully closed, representing the worst-case scenario from a flood perspective.

Flow	UNET elevation	HEC-HMS with gates closed	Difference
(cfs)	(ft)	(ft)	(ft)
1120	770.06	770.07	-0.01
1870	770.69	770.67	0.02
2290	770.96	770.97	-0.01
2550	771.15	771.15	0
3250	771.6	771.6	0
3500	771.77	771.75	0.02
4150	772.2	772.13	0.07

 Table 3. Comparison of dam rating curve from the Flood Insurance Study UNET model and the HEC-HMS model with all dam gates closed.

To evaluate whether the proposed change in the target stage during the period of interest would require a change in the BFE, we used the newly developed HEC-HMS model to evaluate the effect of reduced antecedent storage on the flood stage elevation, assuming in both cases that the dam gates are fully closed.

The antecedent flows in the UNET model output were adjusted to ensure a stable elevation in the impoundment prior to the onset of the flood peak; for the 8.37' lake elevation this required a steady inflow of 351 CFS, while for the 8.5' lake elevation the antecedent flows were increased to 456 cfs. Flows above these values were not changed and were identical in both simulations. The initial lake elevation was set to the target elevation in each model to ensure rapid equilibration.

Under these conditions the HEC-HMS model shows a peak flood stage of 10.82 ft for both scenarios, confirming that the loss of storage from the higher initial stage does not impact the 100-year flood elevation to 0.01 ft, using the conditions represented in the effective Flood Insurance Study model.

#### Actual Impact on Flood Elevations

While the above analysis evaluated the proposed conditions compared with the conditions modeled in the FIS (with a starting lake stage of 8.37 ft), it is also relevant to compare flood elevations for lake stages starting at 8.0 ft vs. 8.5 ft. This comparison applies to the proposed extension of the summer maximum level, from May 1 – May 20 and October 1 – October 15.

With this in mind, we repeated the above analysis, comparing the 100-year flood elevation for starting stages at the winter and summer maxima (8.0 ft and 8.5 ft, respectively). Simulating the 8.5 ft stage requires a 456 cfs antecedent flow (as before), while the 8.0 ft stage is initialized with a flow of 110 cfs. Running the same 100-year hydrograph in each simulation results in a slightly lower peak flood stage of 10.81 ft for the winter level scenario vs 10.82 ft for the summer level scenario.

While this shows an increase in the modeled 100-year flood stage of 0.01 ft, it must be noted that this increase in risk would occur only over the 19 days in May and 14 days in October during which the increase in stage is requested, or less than 10% of the year. It also assumes that the gates are fully closed and are not opened in anticipation of or in reaction to the flood event, which is highly conservative.

We therefore concluded that the proposed operation change would have minimal impact on flood elevations of Buffalo Lake.

#### 4. **GROUNDWATER IMPACTS**

#### 4.1. Issues Addressed

This analysis evaluated how the proposed increased duration of the summer maximum lake stage would affect groundwater near the lake. In particular, impacts to agricultural lands were considered, including tile drainage systems and depth to water table changes. The proposed dam operation changes affect the timing of groundwater fluctuations but not the magnitude of variations. We evaluated the distance from Buffalo Lake at which the seasonal lake level fluctuation has an impact on groundwater elevations, and evaluated the time it takes groundwater to respond to the seasonal change in lake level.

#### 4.2. Areas Evaluated for Risk

#### 4.2.1. Agricultural Areas

Potential risks to agriculture stem from the longer period that is proposed for the summer maximum, which could coincide with spring planting or fall harvest activities. Impacts in fall appear less likely than in spring, because the WDNR's water level records indicate that the lake is commonly below its summer maximum at the end of the summer (before the date when the winter maximum takes effect) due to dry conditions and low lake inflows, however during and after the spring snowmelt it is likely the lake could be brought to the proposed higher target elevation in many years. Low-lying agricultural fields near the lake with the potential to be impacted by increased groundwater elevations were identified based on information from the District and aerial photograph review. Specific areas evaluated are shown in **Figure 6** and described below.

#### Agricultural Area 1

These fields are upstream of Buffalo Lake along the west side of the Fox River near the CTH O crossing. Marquette County LiDAR elevation data indicate that the Fox River elevation here was approximately 2 ft higher than the level of Buffalo Lake during that 2018 survey date, with a river elevation of 772.2 ft. The lowest elevation of crop fields in this area is approximately 780 ft.

#### Agricultural Area 2

These fields east of Endeavor on the east side of Buffalo Lake south and north of Gem Avenue. A ditch network drains these fields to the Fox River upstream of Buffalo Lake. Field elevations range from 772 ft in the southern part of this area to over 780 ft in the northern portion near and to the north of Gem Avenue.

#### Agricultural Area 3

This area is south of Packwaukee on the east side of Buffalo Lake, with fields primarily south of the railroad tracks. Ditches drain these fields southward to the Fox River near the upstream end of Buffalo Lake. Fields in this area are primarily above elevation 775 ft.

#### Agricultural Area 4

This muck farm is located at Endeavor west of Buffalo Lake and Interstate 39. This is a former wetland basin that drained to the Fox River via Chapman Creek. A ditch network is visible throughout the farm, but fields are generally lower than the level of Buffalo Lake and Chapman Creek, with elevations of 763 – 767 ft throughout much of the farm and 769 – 771 at southern fields near the creek. The fields are presumably dewatered by a pumping system because there is no route for gravity drainage, and pumps have been observed by a District representative (Dustin Esselman, written communication, 2025).

#### Agricultural Area 5

Fields north of Buffalo Lake at Packwaukee are drained by the Mad River and tributary ditches. Most of the fields in this area are above elevation 780 ft, but the southern portion of this area has fields at 774 – 775 ft.

#### 4.2.2. Residential Areas

Potential residential impacts include groundwater interference with septic systems and basement seepage. Because the proposal only extends the duration of the summer maximum but does not increase that level, the change would only be expected to prolong existing problems with high groundwater, if they are currently occurring.

Residential areas are present along most of the shoreline of the lake, in Packwaukee and Montello, Buffalo Shore Estates on the east side of the lake between Packwaukee and Endeavor, and on the west side of the lake at Endeavor. These areas include some low-lying properties near the lake where high groundwater could have an impact on septic systems or other features.



Figure 6. Agricultural areas evaluated for groundwater impacts.

#### 4.3. Aquifer Properties

Properties of the groundwater flow system around Buffalo Lake were reviewed to provide insights into potential impacts of the proposed dam operation change and to develop input parameters for groundwater modeling techniques used to quantify groundwater response to lake level changes.

The Buffalo Lake area has not been subject to a detailed hydrogeologic study, however it is at the boundary of the WDNR's Central Sands Lakes Study<sup>1</sup>, and the geology has been mapped by the Wisconsin Geological and Natural History Survey. Additional information is provided by Well Construction Reports available from the WDNR.

Buffalo Lake and the Fox River occupy an area once flooded by Glacial Lake Oshkosh<sup>2</sup>. Up to 80m of clay were deposited in offshore areas of the glacial lake, and sandier deposits formed closer to the lake's shoreline. Sandy glacial till deposits and sandstone bedrock are present below the lake deposits. The geologic map of the area indicates the presence of peat over sandy and silty wetland and stream deposits at the west end of Buffalo Lake (map units po and ps on **Figure 7**). This suggests the possible presence of high-transmissivity sandy materials in that area.

EOR's review of Well Construction Reports for 25 wells near Buffalo Lake found that most wells near the western part of Buffalo Lake (**Figure 8**), where low lying farms are located, were drilled through sand overlying clay and sandy till and/or sandstone bedrock. The shallow sand is typically tens of feet thick at these wells, with a mid-range value of about 50 ft. Many of them have little water level drawdown reported during drillers' pumping tests, indicating high hydraulic conductivity. It therefore appears that the water table around Buffalo Lake is in a sandy aquifer that is likely to be well connected to the lake.

<sup>&</sup>lt;sup>1</sup> Wisconsin Geological and Natural History Survey, 2020. Appendix a – Central Sands Lakes Study Technical Report: Data Collection and Hydrostratigraphy.

<sup>&</sup>lt;sup>2</sup> Hooyer, TS, Mode WN and Clayton, L 2021. Quaternary geology of Columbia, Green Lake, and Marquette Counties, Wisconsin, with contributions to the map by JW Attig: Wisconsin Geological and Natural History Survey Bulletin 114, 38p, 1 pale, scale 1:100,000.



Figure 7. Quaternary geology map of the Buffalo Lake area (from WGNHS).



Figure 8. Locations of Well Construction Reports near west end of Buffalo Lake. Records for 25 of these wells were reviewed for geologic and groundwater information.

The WDNR's Central Sands Lakes Study included 46 aquifer tests for wells in the region, albeit farther north than Buffalo Lake. The mean hydraulic conductivity and specific yield for these tests were 106 ft/d and 0.17, respectively. The WDNR study also used water supply well drillers' specific capacity tests for approximately 23,000 wells in the unconsolidated aquifer. The average hydraulic conductivity for wells east of the terminal glacial moraine, where Buffalo Lake is located, was 112 ft/d (**Figure 9**).



Figure 9. Hydraulic conductivity of the unconsolidated aquifer from well drillers' specific capacity tests (Figure 31b from the WDNR Central Sands Lakes Study)

#### 4.4. Impact Analysis

This analysis evaluated groundwater impacts related to the proposed extension of the summer maximum lake stage period. We used the groundwater GFLOW to evaluate the extent of the area around the lake where groundwater elevation is affected by changes between lake stages 8.0 ft and 8.5 ft. The length of time it takes for groundwater to respond to a change in lake level was evaluated using transient analytical equations.

#### 4.4.1. Extent of Groundwater Affected by Lake Level Change

A groundwater model was constructed using the computer program GFLOW to evaluate differences in the groundwater elevation for lake stages 8.0 ft and 8.5 ft. GFLOW is a steady-state, 2-dimensional analytic element model distributed by the U.S. Environmental Protection Agency that is well-suited to simulate groundwater-surface water interactions. The model simulates the regional flow system using Buffalo Lake, the Fox River and tributary streams as head boundary conditions, a regional recharge rate of 11.5 in/yr estimated by the U.S. Geological Survey<sup>3</sup>, and the properties of the shallow sand aquifer described above (hydraulic conductivity of 112 ft/d; aquifer thickness of approximately 50 ft, and porosity of 0.2).

The groundwater model simulates flow toward Buffalo Lake and tributary streams, as expected. Water table elevation contours for the lake at 8.0 ft and 8.5 ft show similar patterns with little difference visible at a scale that includes the extent of the lake (**Figure 10**). Because the model is steady state, it simulates the maximum impact of the lake level change and does not provide information on how rapidly water table fluctuations occur. The water table also fluctuates with climatic conditions (i.e. wet and dry seasons), which are not simulated by this GFLOW model. The simulations here illustrate the difference that lake stage makes for groundwater levels.

The extent of the area around Buffalo Lake where groundwater elevations differ for the lake at 8.0 ft versus 8.5 ft is shown in **Figure 11**. This distance varies around the lake but is on the order of one mile in many locations. More rise generally occurs where streams drain into the lake, because the water level in streams is affected by the increase in lake level, and this contributes to groundwater rise near the streams. The seasonal water table rise at the agricultural fields described above ranges from less than 0.1 ft to 0.3 ft (**Table 4**), with the exception of Area 4 where pumping is used for dewatering. The GFLOW model does not represent impacts on that area, which are discussed below in Section 4.4.3.

The seasonal water table rise also occurs in residential areas, including the northern shoreline of the lake, Packwaukee, Buffalo Shore Estates, and at Endeavor. We are unaware of high groundwater impacts to septic systems or other uses of residential properties around the lake. Additional impacts due to the proposed changes to the dates for the summer maximum are unlikely because (1) the lake is already managed at the

<sup>&</sup>lt;sup>3</sup> Gebert, WA, JF Walker, and RJ Hunt, 2011. Groundwater Recharge in Wisconsin - Annual Estimates for 1970-99 using Streamflow Data. USGS Fact Sheet 2009-3092.

8.5 ft stage for more than 4 months of the year, and (2) historical information suggests the lake was typically higher than 8.5 ft before the WDNR began to more actively manage the water level in 2019.



Figure 10. Simulated water table elevation contours for lake stages 8.0 ft and 8.5 ft. Contour interval is 2 ft.

Agricultural Area	Simulated rise in water table for lake stage increase from 8.0 ft to 8.5 ft	
Area 1	<0.1 ft	
Area 2	0.2 ft	
Area 3	0.1 – 0.2 ft	
Area 4	N/A <sup>1</sup>	
Area 5	0.1 – 0.3 ft	

Table 4. Seasonal water table rise for lake stage increase from 8.0 ft to 8.5 ft for selected agricultural areas.

<sup>1</sup> Model simulation of water table change does not apply to Area 4.



Figure 11. GFLOW model simulation of rise in water table for lake stage increase from 8.0 ft to 8.5 ft.

#### 4.4.2. Transient Response Time

The purpose of this analysis was to understand how the change in timing of the rise in lake level in spring and drop in lake level in fall would affect properties near the lake with shallow groundwater. This timing is particularly relevant for agricultural fields where spring planting and fall harvest activities could be affected. We used the analytic method described by Kresic<sup>4</sup> (**Figure 12**) to estimate how long it takes the groundwater near the lake to rise or fall in response to a change in lake level. This technique uses the aquifer properties described above to calculate the rate at which groundwater rises in response to the seasonal 0.5 ft lake level increase at different distances from the lake. This method assumes that the change in lake level is rapid compared to changes in the groundwater level, which is reasonable because lake level data from the WDNR indicate that the spring rise in lake level from 8.0 ft to 8.5 ft typically takes about a week.



Figure 12. Boundary conditions for transient 1-dimensional flow with a sudden change at a boundary, such as a lake (from Kresic, 1997)

This analysis simulates the water table response in the shallow sand aquifer connected to the lake, using the same aquifer hydraulic conductivity (112 ft/d) and thickness (50 ft) as the GFLOW model, plus the specific yield value of 0.17 determined by the WDNR Central Sands Lakes Study. The analytical equation was used to compute water table change over time at different distances from the lake that fall within the zone of influence around the lake determined by the GFLOW model.

The calculated water table change rates (**Table 5**, **Figures 13 - 15**) illustrate that the shallow sandy aquifer is transmissive enough that the water table would rise quickly after the lake is elevated 6 inches in May. At a distance of 100 ft, most of the groundwater rise would occur within about a week. At 500 ft, the groundwater would rise 4 inches after 20 days (the proposed number of days to change the start of the summer maximum). At distances of 1000 ft and greater, the water table rise would occur more slowly, taking more than a month to rise 3 inches. This analysis illustrates that near the lake, groundwater response is

<sup>&</sup>lt;sup>4</sup> Kresic, N, 1997. Quantitative Solutions in Hydrogeology and Groundwater Modeling. Lewis Publishers.

likely to be fast enough to make a measurable difference in groundwater levels in May if the lake is raised to the summer maximum level on May 1. Farther from the lake, most of the groundwater rise would occur during the summer, presumably after spring planting activities have been completed.

	Water Table Change (inches)			
Distance from Lake (ft)	5 days	20 days	45 days	
100	5 in	6 in	6 in	
500	2 in	4 in	5 in	
1000	0 in	2 in	3 in	
1500	0 in	1 in	2 in	

 Table 5. Analytical calculations of time for water table elevation change at different distances from Buffalo

 Lake following a 6 inch increase in lake level, rounded to the nearest inch.



Figure 13. Predicted water table rise after the lake is raised from 8.0 ft to 8.5 ft at 100 ft from the shoreline.



Figure 14. Predicted water table rise after the lake is raised from 8.0 ft to 8.5 ft at 500 ft from the shoreline.



Figure 15. Predicted water table rise after the lake is raised from 8.0 ft to 8.5 ft at 1500 ft from the shoreline.

#### 4.4.3. Muck Farm at Endeavor (Area 4)

As noted above, this farm is unlike other areas evaluated because its fields are below the water level of Buffalo Lake and must be dewatered by pumping. At higher the lake levels, there is a greater head difference between the lake and the field drainage system, so that a higher pumping would be needed to maintain the same water level in the fields. The proposed operation change does not increase the summer maximum lake stage, but it would increase the time period when a higher dewatering rate could be necessary. Without site-specific calibration data, the groundwater model can not estimate the increased pumping rate with confidence.

The eastern edge of the fields at CTH CX is approximately 1500 - 2000 ft from Buffalo Lake, and the southern edge of the fields is immediately adjacent to Chapman Creek, which could experience some water level rise with the change in lake level. The analysis above indicates that some difference in groundwater levels at the farm could be experienced in May if the lake is raised to 8.5 ft on May 1.

More information from this farmer would be helpful in determining the magnitude of impact on their operations. For example, does the existing dewatering system have the capacity to get the fields dry enough when the lake is at the 8.5 ft summer maximum? How much does additional dewatering at the higher lake stage cost in terms of additional energy usage? When are the critical periods that fields need to be sufficiently dewatered?

#### 5. THERMAL IMPACTS

#### 5.1. Data Collection

Cason Land & Water Management collected water temperature data at two locations in 2024. The 2024 monitoring began on May 30, 2024 below the dam gates (**Figure 16**) and upstream of the lake at CTH O bridge (**Figure 17**). The logger below the dam gates was removed on September 6, 2024 because no flow was passing through the gates. The logger upstream of the lake was removed on October 15, 2024. Monitoring for 2025 began with deployment of data loggers on April 4; results are not available at the time of this writing.



Figure 16. Temperature monitoring location downstream of the dam in 2024.



Figure 17. Temperature monitoring location upstream of Buffalo Lake in 2024.

#### 5.2. Analysis

Water temperature in the Fox River downstream of the Buffalo Lake dam could be affected by a change in discharge related to the proposed water level change. Discharge downstream of the dam is reduced in the spring during the time that the dam gates are adjusted to hold back more water and raise the lake level from 8.0 ft to 8.5 ft, and flow is increased in the fall if it is necessary to open gates to release more water to lower the lake level. The proposed operation change would shift the spring period of reduced discharge 20 days earlier (starting on May 1 instead of May 20), and the period of increased fall discharge would occur two weeks later (starting on October 15 instead of October 1).

As noted above, the time for the lake to adjust between the winter and summer maxima is approximately 1 week. This means that the Fox River discharge downstream of the dam would be lower than for current operation from about May 1 - 7, and discharge downstream would be greater than for current operation from about May 20 - 27 (the current adjustment period). The magnitude of discharge reduction in the spring has historically been about 150 - 250 cfs, based on WDNR's calculated outflows at the dam before and after gate adjustments.

In fall, release of water from the dam is not always needed to reach the winter maximum stage of 8.0 ft due to low flows and lake levels, as noted above. If a release is necessary, the increase in downstream discharge would be shifted from October 1 to October 15.

Temperature monitoring data do not capture the period when the water level was increased from 8.0 ft to 8.5 ft in 2024, and data loggers near the dam were removed due to lack of flow before the October 1 transition to the winter maximum stage. Data currently being collected in 2025 should capture the period when the lake is raise to the summer maximum. Data from 2024 (**Figure 18**) illustrate that the temperature difference from upstream of the lake to downstream of the lake varied a few degrees around zero from the end of May to mid-June, indicating little difference in the daily average temperature upstream and downstream of the lake. From mid-June through mid-August, the temperature downstream of the lake was about 2 degrees warmer than the upstream temperature, indicating the warming effect of the lake. After mid-August, the data show substantial scatter, likely due to low-flow conditions at both sites.

2025 monitoring data will provide more information, but available data show that the period in early May when discharge in the river downstream of the dam would be lower for the proposed operation change is before the lake began to have a warming effect in 2024. *This observation, the short duration of reduced downstream flows (approximately 1 week), and the fact that early May is not typically a critically cold or hot part of the year suggests that downstream temperature changes will be minimal for the proposed operation change.*


Figure 18. 2024 temperature data and upstream-downstream difference. (Data from Cason Land & Water Management)

# 6. SHORELINE EROSION RISK

#### 6.1. Methods & Data

Increased water depth is a potential risk factor for shoreline erosion, and this risk was evaluated using the WDNR Erosion Intensity Score Worksheet. This is a semi-quantitative tool that considers: fetch; water depth; bank height, composition, stability, vegetation, orientation and geometry; adjacent structures; aquatic vegetation; and boat wakes.

The WDNR tool was applied for the east end of Buffalo Lake where wave erosion risk is highest, because it has the longest fetch relative to prevailing west winds, the deepest water, and least aquatic vegetation cover to dissipate energy. May and October could be sensitive periods; differences in aquatic vegetation growth from early to late May and potential senescence in early-mid-October could result in different aquatic plant cover during the period when the summer maximum lake stage is proposed to be extended.

Erosion Intensity Scores were calculated for existing conditions, representing the lake at 8.0 ft, and for an increase in lake level to 8.5 ft. To be conservative, worst-case rankings were used for the 8.5 ft scenario, increasing water depth at 20 ft and 100 ft offshore by one category and reducing aquatic vegetation cover by 1 category.

Data sources included maps of the lake shoreline, the 1967 bathymetric map, soil survey data, LiDAR topographic data for the shoreline, and observations of aquatic vegetation and shoreline condition visible on aerial photographs.

We calculated "high" Erosion Intensity Scores for both scenarios, with numeric scores of 55 for the 8.0 ft stage and 60 for the 8.5 ft stage.

Because this potential change in erosion potential only applies from May 1 –20 and October 1 – 15 and even conservative representations of the changing conditions results in the same Erosion Intensity Score category, we conclude that the risk of increased shoreline erosion for the proposed operation change is small. The small expected change in shoreline erosion would have a minimal impact on nutrient loading from lakeshore sediment. In fact, a higher water level would be expected to slightly increase sediment and nutrient trapping by particle settling and reduce potential resuspension of lakebed sediments and the nutrients they contain.

# 7. WETLAND IMPACTS

The purpose of this analysis was to evaluate potential effects from proposed water level manipulations on the reservoir, with proposed changes extending high water levels on Buffalo Lake 3 weeks earlier in spring and 2 weeks later in fall. EOR used a desktop data review and a field visit to key areas to evaluate this issue.

#### 7.1. Desktop Data Analysis

#### 7.1.1. Wetland Communities Near Buffalo Lake

According to Natural Heritage Inventory Data, upland natural communities occurring within five miles of Buffalo Lake include Oak Barrens, Southern Dry Forest, Dry Prairie, Northern Dry Forest, Northern Dry-Mesic Forest, Mesic Prairie, Oak Woodland, and Eastern Red Cedar Thicket.

Wetland natural communities within five miles include Northern Wet Forest, Southern Sedge Meadow, Floodplain Forest, Calcareous Fen, Southern Tamarack Swamp, Emergent Marsh, Northern Sedge Meadow, Wet-Mesic Prairie, Shrub-Carr, Spring Pond, Open Bog, Lake (shallow, hard, seepage), Lake (deep, hard, drainage), Lake (shallow, soft seepage), and Springs and Spring Runs—Hard.

#### 7.1.2. Aerial Imagery Review/Offsite Analysis

Three areas were the focus of historic aerial imagery analysis, including a hay field southeast of County Highway K, Page Creek Marsh State Natural Area, and Summerton Bog State Natural and the surrounding muck farms. Twelve photos were obtained from the 1990's to 2022.

The hay field southeast of County Highway K, or Page Creek Southeast, appears to be a hay field as late as 2010, with about half the field harvested for hay and the other half left fallow. By 2013, the area was row-cropped, with minimal crop stress but some slight saturation signatures. Overall, changes in this area of interest appear to be linked to general climatic trends and broad land use changes, rather than lake level manipulation.

Page Creek Marsh State Natural Area shows a shift from open fields in upland areas to dominance by scrubby oaks and invasive brush. By 2022, small clumps of dogwood and willow are visible in the wetland and along its edges. Off-color areas within the wetland indicate invasion of phragmites and cattails, a symptom of invasive species spreading throughout the landscape. Changes here do not appear linked to any apparent water level manipulation.

Summerton Bog shows muck farms in a historic wetland basin, as well as the tamarack swamp present within the State Natural Area. The muck farms appear to have significant dikes and drainageway infrastructure and likely pump water out of the fields. Tamaracks within the State Natural Area appear healthy as late as 2015, but by 2020 there is a significant die off of tamarack.

#### 7.2. Field Observations

A site visit was conducted in the area surrounding Buffalo Lake near Montello, Wisconsin on November 19, 2024. EOR staff looked for signs of stress on woody species that may have included adventitious roots on tamarack, needle drop (needles were still visible on November 19), die-off of wetland trees, or community composition change in calcareous fens.

Soil pits were sampled to a depth of at least 24 inches at Page Creek Marsh State Natural Area and Summerton Bog State Natural Area. Saturation and water table depth were observed. A species list of all vascular plants was taken at Page Creek Marsh. Other qualitative data observed included natural communities present, current water levels, geomorphology of the reservoir, and stress to woody plants including adventitious roots or needle-drop on coniferous trees.

Water table depth, where reached, ranged from 0-16 inches, with the water typically at or very near the surface in wetlands. Antecedent precipitation conditions were normal for the preceding three months, with the last month receiving above average precipitation (**Figure 20**). At Page Creek Marsh, the water table corresponded tightly with elevation. At Summerton Bog, the water table occurred at 4-6 inches along a seepage slope, with springs and spring runs emanating along this slope. Below the seepage slope, topography leveled off into a tamarack swamp, and the water table occurred at 12 inches.

Elevation	Water Table Depth	Site
794 ft	4-6 in	Summerton Bog
783 ft	12 in	Summerton Bog
772 ft	1 in	Page Creek
771 ft	+1 in (standing water)	Page Creek
771 ft	+2 in (standing water)	Page Creek
773 ft	11 in	Page Creek
773.5 ft	15 in	Page Creek

#### Table 6. Water table observed at soil pits.

Groundwater influences both the Sedge Meadow (Page Creek) and Tamarack Swamp (Summerton Bog), with fen species such as *Muhlenbergia glomerata*, *Rumex Britannica*, and *Sium suave*, indicating contact with calcium-rich groundwater. Uplands in topographically higher positions contained classic Oak Barrens species like *Quercus ellipsoidalis*, *Schizachyrium scoparium*, and *Asclepias verticillata*. More disturbed areas were dominated by non-native shrubs and *Bromus inermis*. A full species list from Page Creek Marsh is located in **Appendix B**.

At Page Creek Marsh, shrub species are not common in the wetland, although *Cornus racemosa* is located on the toeslope of the wetland with a mix of Facultative species. Aerial imagery analysis indicates that this shrub invasion is more recent, occurring within the last ten years, although this is likely the result of fire suppression rather than impacts from lake levels. The WDNR's <u>Prescribed Burn Dashboard</u> indicates that no fires have occurred at Page Creek Marsh since at least 2019.

Numerous dead tamarack trees were observed at Summerton Bog and along Lakeview Drive, though no adventitious roots were observed, indicating a lack of stress from flooding.<sup>5</sup> The dozens of dead tamaracks observed on site near the soil pits did not show adventitious roots, but there is considerably more die off to the south. Die off could be resulting from larch beetle or larch sawfly, from high water levels and flooding in 2018, or from lower water levels due to active operation of the dam gates since 2019. Evidence collected onsite is inconclusive about the cause of tamarack die off.

<sup>&</sup>lt;sup>5</sup> Calvo-Polanco M, Señorans J, Zwiazek JJ. Role of adventitious roots in water relations of tamarack (Larix laricina) seedlings exposed to flooding. BMC Plant Biol. 2012 Jun 27;12:99. doi: 10.1186/1471-2229-12-99. PMID: 22738296; PMCID: PMC3431261.

A potential White Pine Swamp was observed along Highway K, south of Buffalo Lake, although this community has little apparent connection to Buffalo Lake, and with drainage in this area southwest to Williams Lake.



Figure 19. Bordering Page Creek



Figure 20. Antecedent precipitation for the Buffalo Lake area

#### 7.3. Interpretation and Conclusions

Effects from manipulating lake levels six inches in spring and fall are unlikely to impact wetland hydrology, native plant communities, or rare/protected resources. Natural communities likely to be sensitive to such disturbances include groundwater and lake level-dependent communities such as calcareous fens or emergent, floating-leaved, or submergent marsh, although evidence collected from the site visit indicates that impacts are unlikely.

Within preserved natural areas, groundwater resources appear to be intact, with observations from the site visit showing evidence of groundwater at or near the surface of both Page Creek and Summerton Bog. Emergent, floating-leaved, or submergent marsh along Buffalo Lake are also unlikely to be impacted since most of the lake is shallow enough to support aquatic macrophytes. Maintaining higher water levels may impact plant germination, though this is unlikely to have any significant effect since drawdowns will still occur in winter, allowing for seed contact with nearshore habitat.

Based on data gathered during the site visit, wetland natural communities are intact, with healthy hydrologic profiles. There is no evidence of impacts from either flooding or drawdowns. Minor indications of degradation include invasions of shrubs and invasive species. Based on aerial imagery review, these invasions had occurred since at least 2005, with shrub invasion occurring in the last ten years. Both symptoms are indicative of widespread changes occurring elsewhere on the landscape such as habitat fragmentation, sprawling home developments, and the spread of invasive species. A comparison of Mean Coefficient of Conservatism values shows that the Page Creek Sedge Meadow retains a high conservation value (**Table 7**). The Mean C value was lower on the November 19 site visit due to the senescence of most plant species. *Carex* species dominate the Sedge Meadow, and since most Carex were unidentifiable on the November site visit, the 2024 Mean C value is an underestimate.

Survey	Cover-weighted Mean C
WDNR Timed Meander (7/7/2015)	5.828
EOR Meander (11/19/2024)	4.780

Table 7. Cover-weighted Mean C values at Page Creek Marsh State Natural Area

With the steep slopes surrounding most of Buffalo Lake, groundwater-dependent systems are unlikely to be affected by the maintenance of high lake levels and may even benefit from higher groundwater levels during the growing season. A comparison to reference Sedge Meadows shows that Page Creek is within the hydrologic profile of these reference wetlands (**Figure 21**). Lake levels appear unlikely to affect native plant community composition and structure, except perhaps immediately adjacent to Buffalo Lake.

The potential White Pine Swamp is unlikely to be affected by manipulated water levels as it occurs far above lake levels.

Overall, current community composition and hydrologic profiles appear healthy, and extending high water levels six weeks into spring and fall is unlikely to affect wetlands or natural communities.

Other potential unknowns include the cause of tamarack die off and potential effects on wetlands adjacent to the lake that were not part of this investigation. There is some apparent tamarack die off along the Interstate on the west side of the Buffalo Lake. The dead trees rapidly increase after 2018, rather than dying all at once. If sudden water level shifts were the cause, a sudden die off would also be expected. Rather, aerial imagery suggests this is a gradual die off which might be consistent with larch beetle or larch sawfly infestations. If cause and effect are desired, a more focused study of tamaracks should be investigated.

Wetlands absent from this investigation could be affected by manipulated lake levels, but Page Creek Marsh and Summerton Bog are some of the closest and highest quality wetlands near Buffalo Lake. The most sensitive systems are expected to be groundwater dependent, and both communities analyzed showed healthy groundwater profiles. It is expected that this is the case in other wetlands surrounding Buffalo Lake, but further monitoring could take place to confirm this assumption.

Overall, keeping lake levels raised for 5 more weeks is unlikely to negatively impact these wetland communities and is more likely to provide stable conditions throughout the growing season.





Figure 21. Plots from well data at WDNR reference sites, with Page Creek Marsh water levels also shown (red star).

## 8. OVERALL CONCLUSIONS

#### Shoreline fluctuation

 No detailed elevation data is available to precisely delineate the shoreline of the lake at stages of 8.0 ft and 8.5 ft, or even 9.0 ft; however, historic aerial photographs from periods when the lake stage was measured show that lake inundation extent and surface area do not vary much in the range of the winter and summer maxima.

#### Lake Flood Levels

- The dam gates have the capacity to substantially lower lake flood stage.
- The proposed operation change will not impact the Base Flood Elevation of Buffalo Lake defined in the Marquette County Flood Insurance Study.
- The difference in 100-year lake stage elevation is minimal for a flood that occurs when the lake starts at 8.0 ft vs. 8.5 ft.

#### Groundwater

- Groundwater levels around Buffalo Lake rise and fall depending on the lake stage. The zone near the lake affected by changes between the winter and summer maxima is variable and extends up to about 1 mile in some locations.
- The change in groundwater level in this zone of influence is rapid enough to create measurable changes in groundwater levels during the proposed extension of the summer maximum stage in spring and fall.
- Additional risk to residential properties appears minimal, but existing problems with high groundwater (if they are occurring) could be extended with an increased duration of the summer maximum lake stage.
- The predicted rise in groundwater level due to increasing the lake stage from 8.0 ft to 8.5 ft evaluated at 4 farms to the south, east and north of Buffalo Lake ranges from less than 0.1 ft to 0.3 ft.
- The muck farm west of the lake at Endeavor appears to have a dewatering pumping system that would be affected by an extended duration of the summer maximum lake level. A higher lake level would require a higher dewatering rate; the high stage is already occurring each summer, but the duration requiring more pumping would be extended by 20 days in May.
- We recommend that the District engages with this muck farmer to discuss the proposed operation change, potential impacts on farm operations, and options for mitigation, if necessary.

#### Fox River Temperature

- The time to raise the lake from the winter level to the summer level has historically been about one week, and during this period, flows downstream in the Fox River are affected.
- This period of approximately 1 week of reduced spring flows would shift from late May 20 to early May with the proposed operational change.
- Lake level data are insufficient to demonstrate the historical time to draw the lake down in the fall, because fall water levels have commonly started below the summer maximum due to dry weather. However, the capacity of the gates suggests a similar duration for this lake level adjustment as for spring.
- In May of 2024, Buffalo Lake had little impact on water temperature downstream.
- The proposed operation change is expected to have minimal impact on the temperature of the Fox River downstream of the lake, given the short duration of flow reduction and typical lack of thermal stress (extreme hot or cold) in May. In fact, the shift in discharge reduction from late May to early May could reduce warm weather thermal stress in the river.

#### Shoreline Erosion

- The shoreline Erosion Intensity Score is high for the lake at both the winter and summer max. Changes in shoreline erosion potential due to the proposed operation change are minimal.
- This implies minimal change in nutrient loading to the lake from shoreline erosion.
- At the higher lake level, a slight increase sediment and nutrient trapping by particle settling and reduced resuspension of lakebed sediments and the nutrients they contain is expected.

#### Wetlands

- Wetlands evaluated with desktop data and field visits do not show indications of stress related to lake level fluctuations.
- Extending the duration of the summer maximum lake level is not expected to negatively affect wetlands and is more likely to benefit them through more stable groundwater levels during the growing season.

# **APPENDIX A. SELECTED SHORELINE AERIAL PHOTOGRAPHS**

The following images were obtained from Google Earth. The dates shown are the dates listed by Google Earth but we are unable to vouch for their accuracy. Locations were selected to represent areas where area change would be visible. While lake stages were not available on the day in question, the recorded stages immediately before and after the photo date are shown. The range of elevations shown here exceeds the range of the proposed changes to the target elevation, and there is little evidence that this change leads to a significant change in the lake surface area.

	March 11 2021	June 30 2021	May 4 2024
	Stage 8.14' (3/10/21)	Stage 8.64' (6/28/21)	Stage 8.03' (4/30/24)
	Stage 8.22' (3/12/21)	Stage 8.5' (7/01/21)	Stage 7.93' (5/07/24)
Boat Landing, Hwy C and 10 <sup>th</sup> Dr		the second	
County D and Lake St			
		·	·



Scientific Name	Family	Native?	С	W	Physiognomy	Common Name
Alnus incana	Betulaceae	native	4	-3	shrub	mountain alder
Asclepias incarnata	Apocynaceae	native	5	-5	forb	swamp milkweed
Bidens tripartita	Asteraceae	native	5	-3	forb	straw-stem beggar-ticks
Bromus kalmii	Poaceae	native	8	0	grass	arctic brome
Calamagrostis canadensis	Poaceae	native	5	-5	grass	blue-joint grass
Carex hystericina	Cyperaceae	native	3	-5	sedge	bottlebrush sedge
Carex lacustris	Cyperaceae	native	6	-5	sedge	common lake sedge
Cornus sericea	Cornaceae	native	3	-3	shrub	red osier dogwood
Dulichium arundinaceum	Cyperaceae	native	9	-5	sedge	pond sedge
Euthamia graminifolia	Asteraceae	native	4	0	forb	common flat-topped goldenrod
Eutrochium maculatum	Asteraceae	native	4	-5	forb	spotted joe-pye-weed
Geum macrophyllum	Rosaceae	native	6	-3	forb	big-leaved avens
Juncus tenuis	Juncaceae	native	1	0	forb	path rush
Lathyrus palustris	Fabaceae	native	5	-3	forb	marsh pea
Lycopus americanus	Lamiaceae	native	4	-5	forb	american water- horehound
Lysimachia thyrsiflora	Primulaceae	native	7	-5	forb	swamp loosestrife
Muhlenbergia glomerata	Poaceae	native	9	-5	grass	marsh muhly
Muhlenbergia mexicana	Poaceae	native	4	-3	grass	leafy satin grass
Onoclea sensibilis	Dryopteridaceae	native	5	-3	fern	sensitive fern

# APPENDIX B. OBSERVED SPECIES LIST, PAGE CREEK MARSH

Rumex britannica	Polygonaceae	native	8	-5	forb	greater water dock
Schoenoplectus						
tabernaemontanı	Cyperaceae	native	4	-5	sedge	great bulrush
Scirpus						
atrovirens	Cyperaceae	native	3	-5	sedge	black bulrush
Scirpus cyperinus	Cyperaceae	native	4	-5	sedge	wool-grass
Sium suave	Apiaceae	native	5	-5	forb	hemlock water-parsnip
Solidago						
canadensis	Asteraceae	native	1	3	forb	canadian goldenrod
Spartina						
pectinata	Poaceae	native	5	-3	grass	prairie cord grass
Spiraea						
tomentosa	Rosaceae	native	6	-3	shrub	hard-hack
Stachys palustris	Lamiaceae	native	5	-5	forb	hedge-nettle
Symphyotrichum						
puniceum	Asteraceae	native	5	-5	forb	swamp aster
Thelypteris						
palustris	Thelypteridaceae	native	7	-3	fern	eastern marsh fern
Typha latifolia	Typhaceae	native	1	-5	forb	broad-leaved cat-tail
Verbena hastata	Verbenaceae	native	3	-3	forb	blue vervain
Zanthoxylum						
americanum	Rutaceae	native	3	3	shrub	common prickly-ash

# **SECTION II**

# Buffalo Lake Enhancement Project Point-Intercept Survey

(Conducted by Cason Land & Water Management, LLC.)



# Buffalo Lake

# 2024 Point-Intercept Survey Summary Report with Statistical Analysis

### Submergent Aquatic Plant Survey

Cason Land & Water Management, LLC conducted a Point Intercept Aquatic plant survey of the Buffalo Lake on July 3<sup>rd</sup> – 26<sup>th</sup>, 2024. At 828 of the 907 grid points (**Figure 1**) plotted across the lake aquatic plant samples were collected from a boat with a single rake pull or throw. At depths of 15 feet or less, a double rake head attached to a pole was used to collect a sample; a double rake head on a rope was used for depths greater than 15 feet. Plants were observed up to a depth of 8 feet (**Figure 2**). All plant samples collected were identified to genus and species whenever possible, and the information was recorded. Twenty-three different aquatic plant species were observed on the rake during the survey and a total of forty-two plant species were observed in total during the survey (**Table 1**). The aquatic invasive species Eurasian watermilfoil and Curly-leaf pondweed were observed during the survey as well as several wetland invasive species which are denoted in red text (**Table 1**). An abundance rating was also given for each species collected using criteria established by the WDNR. In addition to the plant data, water depths were also recorded for each location. Data collected was used to determine species composition, percent frequency and relative abundance.

#### Simpson Diversity Index

To estimate the diversity of the aquatic plant community, the Simpson Diversity Index takes into account both the number of species identified (richness) and the distribution or relative abundance of each species. With the Simpson Diversity Index (D), 1 represents infinite diversity and 0 represents no diversity. That is, the bigger the value of D, the higher the diversity. Buffalo Lake was calculated to have a Simpson Diversity Index of 0.88.

## **Assessment of Floristic Quality Resources**

The plant data collected for Buffalo Lake was used to assess the *floristic quality* of the lake. The method used, assigns a value to each native plant species called a *Coefficient of Conservatism*. Coefficient values range from 0-10 and reflect a particular species' likelihood of occurring in a relatively undisturbed landscape. Species with low coefficient values, such as sago pondweed (*Stuckenia pectinata*) (C=3), are likely to be found in a variety of habitat types and can tolerate high levels of human disturbance. On the other hand, species with higher coefficient values, such as white-stem pondweed (*Potamogeton praelongus*) (C=8), are much more likely to be restricted to high quality natural areas. By averaging the coefficient values available for the submergent and emergent species found in the lake, a value was assigned to the lake. The average *Coefficient of Conservatism* value for lakes in Wisconsin is 6.0, Buffalo Lake's average was also found to be exactly 6.0 during the 2024 survey.

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By utilizing the Coefficients of Conservatism for the plant species of Buffalo Lake, further assessment of floristic quality was made. By multiplying the average coefficient values for Buffalo Lake by the square root of the number of plant species found, a Floristic Quality Index (FQI) was calculated. The average for Wisconsin lakes is 22.2; Buffalo Lake has a FQI of 26.83. According to the U.S. Fish and Wildlife Service "The FQI is an indication of native vegetative quality for an area: generally, 1-19 indicates low vegetative quality; 20-35 indicates high vegetative quality and above 35 indicates "Natural Area" quality. Wetlands with a FQI of 20 or greater are considered high quality aquatic resources."



Figure 1. Point-Intercept survey grid provided by WDNR.

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Figure 2. Maximum Depth of Plant Colonization observed during the 2024 survey.

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#### Table 1. Buffalo Lake Aquatic Species present during the 2024 survey.

Species	Scientific Name	Plant type: floating leaf, free floating, submergent, emergent	% Relative Frequency of Occurence	Sites Found
Coontail	Ceratophyllum demersum	Submergent	17.2	453
Common watermeal	Common watermeal	Free Floating	13.9	367
Small duckweed	Lemna minor	Free Floating	13.8	364
Flat-stem pondweed	Potamogeton zosteriformis	Submergent	12.9	340
Forked duckweed	Lamna triscula	Free Floating	10.3	270
Common waterweed	Elodea canadensis	Submergent	9.8	259
Large duckweed	Spirodela polyrhiza	Free Floating	7.9	208
Eurasian water milfoil	Myriophyllum spicatum	Submergent	5.2	137
White water lily	Nymphaea odorata	Floating leaf	2.4	63
White water crowfoot	Ranunculus aquatilis	Submergent	2	53
Wild celery	Vallisneria americana	Submergent	1.4	36
Southern naiad	Najas guadalupensis	Submergent	0.9	25
Spiny hornwort	Ceratophyllum echinatum	Submergent	0.5	13
Sago pondweed	Stuckenia pectinata	Submergent	0.4	11
Nitella	Nitella sp.	Submergent	0.3	8
Slender waterweed	Elodea nutalli	Submergent	0.2	4
Water star-grass	Heteranthera dubia	Submergent	0.2	5
Curly-leaf pondweed	Potamogeton crispus	Submergent	0.2	6
Leafy pondweed	Potamogeton foliosus	Submergent	0.2	4
Small pondweed	Potamogeton pusillus	Submergent	0.2	5
Muskgrasses	Chara sp.	Submergent	0	1
Northern blue flag	Iris versicolor	Emergent	0	1
Long-leaf pondweed	Potamogeton nodosus	Submergent	0	1
Swamp Milkweed	Asclepias incarnata	Emergent	Visual	1
Wild calla	Calla palustris	Emergent	Visual	1
Bulbet-Bearing Water Hemlocl	Cicuta bilbifera	Emergent	Visual	1
Water Hemlock	Cicuta douglasii	Emergent	Visual	1
Red-osier Dogwood	Cornus sericea	Shrub	Visual	1
Orange jewelweed	Impatiens capensis	Emergent	Visual	2
Purple Loosestrife	Lythrum salicaria	Emergent	Visual	6
American Lotus	Nelumbo lutea	Floating leaf	Visual	11
Ditch Stonecrop	Penthorum sedoides	Forb	Visual	1
Reed canary grass	Phalaris arundinacea	Emergent	Visual	2
Common reed	Phragmites australis	Emergent	Visual	5
Water smartweed	Polygonum amphibium	Emergent/Floating leaf	Visual	3
Great Water Dock	Rumex britannica	Emergent	Visual	1
Common arrowhead	Sagittaria latifolia	Emergent	Visual	1
Willow	Salex sp.	Woody plant	Visual	1
Softstem bulrush	Schoenoplectus tabernaemontani	Emergent	Visual	2
Common bur-reed	Sparganium eurycarpum	Emergent	Visual	9
Cattail	Typha sp.	Emergent	Visual	77
Wild rice	Zizania sp.	Emergent	Visual	18
Filamentous algae	various	Free floating	N/A	63

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Species Richness: 23 Species Richness (with visuals): 42 Simpson Diversity Index (D): 0.88 Floristic Quality Index (FQI): 26.83 Avg. Coefficient of Conservatism (C): 6

The following maps illustrate the distribution of aquatic invasive species (**Figures 3 & 4**), overall rake fullness (**Figure 5**), the seven most abundant (non-free-floating) plant species in Buffalo Lake (**Figures 6-12**), and lastly four floating-leaf species that would have otherwise made the top seven native species list (**Figures 13-16**).

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Figure 3. Distribution of Eurasian Watermilfoil in Buffalo Lake.



Figure 4. Distribution of Curly-leaf pondweed in Buffalo Lake.



Figure 5. Distribution of Total Rake Fullness in Buffalo Lake.



Figure 6. Distribution of Coontail in Buffalo Lake.



Figure 7. Distribution of Flat-stem pondweed in Buffalo Lake.



Figure 8. Distribution of Common waterweed in Buffalo Lake.



Figure 9. Distribution of White water lily in Buffalo Lake.



Figure 10. Distribution of White water crowfoot in Buffalo Lake.



Figure 11. Distribution of Wild celery in Buffalo Lake.



Figure 12. Distribution of Southern Naiad in Buffalo Lake.



Figure 13. Distribution of Common watermeal in Buffalo Lake.



Figure 14. Distribution of Small duckweed in Buffalo Lake.



Figure 15. Distribution of Forked duckweed in Buffalo Lake.



Figure 16. Distribution of Large duckweed in Buffalo Lake.



#### **Trends in the Aquatic Plant Community**

The relative frequency of occurrence of each species found on Buffalo Lake are listed for each year that Point-Intercept surveys were conducted (2015, 2024, **Table 2**). Differences in species richness from year to year are likely due to variable observer biases and accessibility barriers to certain portions of the lake. A year-by-year summary is provided for the number of PI points which were sampled, the species richness observed via rake pulls, species richness including visual observations, the Simpson Diversity index (D) value, the Floristic Quality Index value (FQI), as well as the average coefficient of conservatism (C) value (**Table 3**). However, the FQI and C values were not available from the 2015 survey data.

Statistical analyses were performed on the relative frequency of occurrence data from 2015 to 2024. A chi-square analysis was used to identify statistically significant differences among species and indicate both significant increases and decreases (**Table 4**). As for aquatic invasive species, Eurasian watermilfoil has significantly increased in relative frequency of occurrence whereas Curly-leaf pondweed has significantly decreased in Buffalo Lake (**Table 4**). Additionally for AIS, Brittle Naiad was not observed during the 2024 survey resulting in a significant decrease of that species as well (**Table 4**). Native aquatic plant species that experienced an increase in relative frequency of occurrence during the last nine years include: Common watermeal, Small duckweed, Flat-stem pondweed, Forked duckweed, Large duckweed, Southern naiad, Spiny hornwort, Nitella, American Lotus, Common bur-reed, Cattail, and Wild rice (**Table 4**). Native species that experienced a decrease in relative frequency of occurrence during the last nine years include: Coontail, Common waterweed, White water crowfoot, Wild celery, Sago pondweed, Water star-grass, Small pondweed, Northern watermilfoil, Slender naiad (**Table 4**). Also, three new wetland invasive species were observed during the 2024 survey which were not observed during the 2025 survey, those species include: Purple loosestrife, Reed canary grass, and Common reed.

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Species	Scientific Name	% Relative Frequency of Occurence 2024	% Relative Frequency of Occurence 2015
Coontail	Ceratophyllum demersum	17.2	26.4
Common watermeal	Wolffia columbiana	13.9	2.8
Small duckweed	Lemna minor	13.8	3.2
Flat-stem pondweed	Potamogeton zosteriformis	12.9	5.2
Forked duckweed	Lemna triscula	10.3	0.9
Common waterweed	Elodea canadensis	9.8	19.1
Large duckweed	Spirodela polyrhiza	7.9	0.7
Eurasian water milfoil	Myriophyllum spicatum	5.2	5.1
White water lily	Nymphaea odorata	2.4	2.8
White water crowfoot	Ranunculus aquatilis	2	5
Wild celery	Vallisneria americana	1.4	8
Southern naiad	Najas guadalupensis	0.9	0.2
Spiny hornwort	Ceratophyllum echinatum	0.5	Absent
Sago pondweed	Stuckenia pectinata	0.4	2.1
Nitella	Nitella sp.	0.3	Absent
Slender waterweed	Elodea nutalli	0.2	0.1
Water star-grass	Heteranthera dubia	0.2	3.1
Curly-leaf pondweed	Potamogeton crispus	0.2	2.4
Leafy pondweed	Potamogeton foliosus	0.2	Absent
Small pondweed	Potamogeton pusillus	0.2	8.1
Muskgrasses	Chara sp.	0	0.3
Northern blue flag	Iris versicolor	0	Absent
Long-leaf pondweed	Potamogeton nodosus	0	0.2
Swamp Milkweed	Asclepias incarnata	Visual	Absent
Wild calla	Calla palustris	Visual	Absent
Bulbet-Bearing Water Her	Cicuta bilbifera	Visual	Absent
Water Hemlock	Cicuta douglasii	Visual	Absent
Red-osier Dogwood	Cornus sericea	Visual	Absent
Orange jewelweed	Impatiens capensis	Visual	Absent
Purple Loosestrife	Lythrum salicaria	Visual	Absent
American Lotus	Nelumbo lutea	Visual	0.1
Ditch Stonecrop	Penthorum sedoides	Visual	Absent
Reed canary grass	Phalaris arundinacea	Visual	Absent
Common reed	Phragmites australis	Visual	Absent

+ 6 • . . -Table 2

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#### Table 2 (cont.). Buffalo Lake aquatic plant species present by year and relative frequency of occurrence.

Species	Scientific Name	% Relative Frequency of Occurence 2024	% Relative Frequency of Occurence 2015
Water smartweed	Polygonum amphibium	Visual	Absent
Great Water Dock	Rumex britannica	Visual	Absent
Common arrowhead	Sagittaria latifolia	Visual	Absent
Willow	Salex sp.	Visual	Absent
Softstem bulrush	Schoenoplectus tabernaemo	Visual	Absent
Common bur-reed	Sparganium eurycarpum	Visual	Absent
Cattail	Typha sp.	Visual	0.1
Wild rice	Zizania sp.	Visual	0.1
Filamentous algae	various	N/A	2.7
Northern watermilfoil	Myriophyllum sibiricum	Absent	0.3
Slender naiad	Najas flexilis	Absent	2.8
White-stem pondweed	Potamogeton praelongus	Absent	0.1
Fern pondweed	Potamogeton robbinsii	Absent	0.1
Hardstem bulrush	Schoenoplectus acutus	Absent	0.1
Brittle Naiad	Najas minor	Absent	0.7

Table 3: A year-by-year summary of the number of PI points which were sampled, the species richness observed via rake pulls, species richness including visual observations, the Simpson Diversity index (D) value, the Floristic Quality Index value (FQI), as well as the average coefficient of conservatism (C) value.

	2024	2015
Number of sampled points:	828	675
Species Richness:	23	23
Species Richness (with visuals):	42	23
Simpson Diversity Index (D):	0.88	0.87
Floristic Quality Index (FQI):	26.83	Unavailable
Avg. Coefficient of Conservatism (C):	6.00	Unavailable

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Table 4: Results of chi-square analysis of percent frequency of occurrence survey data from the 2015 and 2024 Point-Intercept surveys of Buffalo Lake. Species are organized by 2024 percent frequency, with the highest frequency first. Green rows indicate significant increase and red rows indicate significant decrease in plant occurrence from 2015-2024. Invasive species are indicated with red text.

Species	Scientific Name	Percent Frequency		Significant	Increase (I) or
Species		2015	2024	Change	Decrease (D)
Coontail	Ceratophyllum demersum	26.4	17.2	**	D
Common watermeal	Wolffia columbiana	2.8	13.9	***	I
Small duckweed	Lemna minor	3.2	13.8	***	I
Flat-stem pondweed	Potamogeton zosteriformis	5.2	12.9	***	I
Forked duckweed	Lemna triscula	0.9	10.3	***	
Common waterweed	Elodea canadensis	19.1	9.8	***	D
Large duckweed	Spirodela polyrhiza	0.7	7.9	***	I
Eurasian water milfoil	Myriophyllum spicatum	5.1	5.2	*	l I
White water lily	Nymphaea odorata	2.8	2.4	n.s.	I
White water crowfoot	Ranunculus aquatilis	5	2	***	D
Wild celery	Vallisneria americana	8	1.4	***	D
Southern naiad	Najas guadalupensis	0.2	0.9	***	I
Spiny hornwort	Ceratophyllum echinatum	0	0.5	**	I
Sago pondweed	Stuckenia pectinata	2.1	0.4	***	D
Nitella	Nitella sp.	0	0.3	*	I
Slender waterweed	Elodea nutalli	0.1	0.2	n.s.	I
Water star-grass	Heteranthera dubia	3.1	0.2	***	D
Curly-leaf pondweed	Potamogeton crispus	2.4	0.2	***	D
Leafy pondweed	Potamogeton foliosus	0	0.2	n.s.	I
Small pondweed	Potamogeton pusillus	8.1	0.2	***	D
Muskgrasses	Chara sp.	0.3	0	n.s.	D
Northern blue flag	Iris versicolor	0	0	n.s.	I
Long-leaf pondweed	Potamogeton nodosus	0.2	0	n.s.	D
Swamp Milkweed	Asclepias incarnata	0	0	n.s.	I
Wild calla	Calla palustris	0	0	n.s.	I
Bulbet-Bearing Water Hemlock	Cicuta bilbifera	0	0	n.s.	I
Water Hemlock	Cicuta douglasii	0	0	n.s.	I
Red-osier Dogwood	Cornus sericea	0	0	n.s.	I
Orange jewelweed	Impatiens capensis	0	0	n.s.	I
Purple Loosestrife	Lythrum salicaria	0	0	*	I
American Lotus	Nelumbo lutea	0.1	0	*	I
Ditch Stonecrop	Penthorum sedoides	0	0	n.s.	1
* significant change ( $\alpha$ =0.05), ** more significant change ( $\alpha$ =0.01), *** most significant change ( $\alpha$ =0.001)					

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Table 4 (cont.): Results of chi-square analysis of percent frequency of occurrence survey data from the 2015 and 2024 Point-Intercept surveys of Buffalo Lake. Species are organized by 2024 percent frequency, with the highest frequency first. Green rows indicate significant increase and red rows indicate significant decrease in plant occurrence from 2015-2024. Invasive species are indicated with red text.

Species Scientific Name	Percent Frequency		Significant	Increase (I) or	
Species	Scientine Name	2015	2024	Change	Decrease (D)
Reed canary grass	Phalaris arundinacea	0	0	n.s.	I
Common reed	Phragmites australis	0	0	*	I
Water smartweed	Polygonum amphibium	0	0	n.s.	I
Great Water Dock	Rumex britannica	0	0	n.s.	Ι
Common arrowhead	Sagittaria latifolia	0	0	n.s.	
Willow	Salex sp.	0	0	n.s.	
Softstem bulrush	Schoenoplectus tabernaemo	0	0	n.s.	I
Common bur-reed	Sparganium eurycarpum	0	0	**	I
Cattail	Typha sp.	0.1	0	***	I
Wild rice	Zizania sp.	0.1	0	**	I
Filamentous algae	various	2.7	N/A	n.s.	I
Northern watermilfoil	Myriophyllum sibiricum	0.3	0	*	D
Slender naiad	Najas flexilis	2.8	0	***	D
White-stem pondweed	Potamogeton praelongus	0.1	0	n.s.	D
Fern pondweed	Potamogeton robbinsii	0.1	0	n.s.	D
Hardstem bulrush	Schoenoplectus acutus	0.1	0	n.s.	D
Brittle Naiad	Najas minor	0.7	0	***	D
* significant change ( $\alpha$ =0.05), ** more significant change ( $\alpha$ =0.01), *** most significant change ( $\alpha$ =0.001)					

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## **SECTION III**

# **Buffalo Lake Enhancement Project** Floating Leaf & Emergent Plant Survey

(Conducted by Cason Land & Water Management, LLC.)



### Buffalo Lake 2024 Floating leaf & Emergent Plant Survey Summary

On August 29, 2024, Cason Land & Water Management, LLC conducted a floating leaf and emergent plant survey on Buffalo Lake which followed a full point-intercept survey.

This survey was conducted by navigating through all navigable waters on Buffalo Lake to plot the diversity and distribution of various floating-leaf, emergent, and riparian plant species in the system. This survey is conducted by trained biologists with experience in plant identification. Locations of these plant communities were delineated using ArcGIS Pro to create the detailed maps provided.

Buffalo Lake experienced higher than average water levels for much of the spring and summer, often over twelve inches higher levels. Based on visual results of the 2024 plant distribution, in comparison with historical aerial imagery and past maps, it appeared that the high water substantially hindered the growth of some species, notably the white-water lily. These comparisons can be clearly made by looking at the attached maps and noting where these plants were historically in comparison to where our polygons were drawn for 2024 observations. Species observed are shown in the table below, with red text indicating invasive species. The distribution of these species is shown in the maps that follow.

Floating leaf & Emergent Plant Species		
American Lotus (Nelumbo lutea)	Orange Jewelweed (Impatiens capensis)	
Arrowhead (Sagittaria sp.)	Phragmites (Phragmites australis)	
Cattail (Typha sp.)	Purple Loosestrife (Lythrum salicaria)	
Blue flag Iris (Iris versicolor)	Softstem Bulrush (Schoenoplectus tabernaemontani)	
Bristly Sedge (Carex comosa)	White Water-Lily (Nymphaea alba)	
Bur Reed (Sparganium sp.)	Wild Rice (Zizania palustris)	
Japanese Knotweed (Reynoutria japonica)	Yellow Iris (Iris pseudacorus)	
Long-leaf Pondweed (Potamogeton nodosus)		

The following deliverables are attached to this survey summary:

• Floating leaf & Emergent Plant Distribution Maps (12).

Thank you for your business and for allowing us to perform this service for your district. If you have any questions, please contact me at 920-290-6810 or LancePaden@CasonLandWater.com

Sincerely, Lance Paden

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